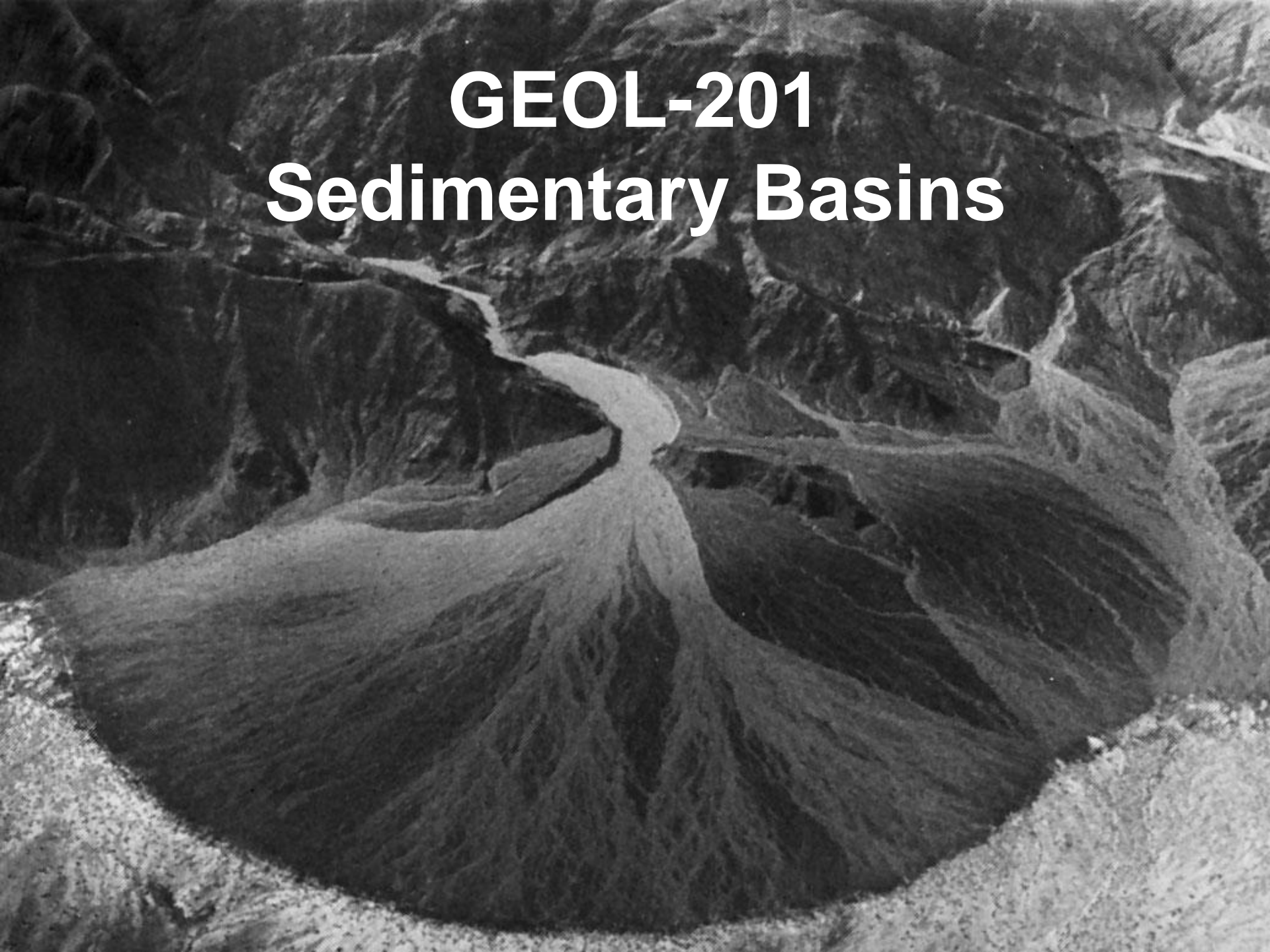
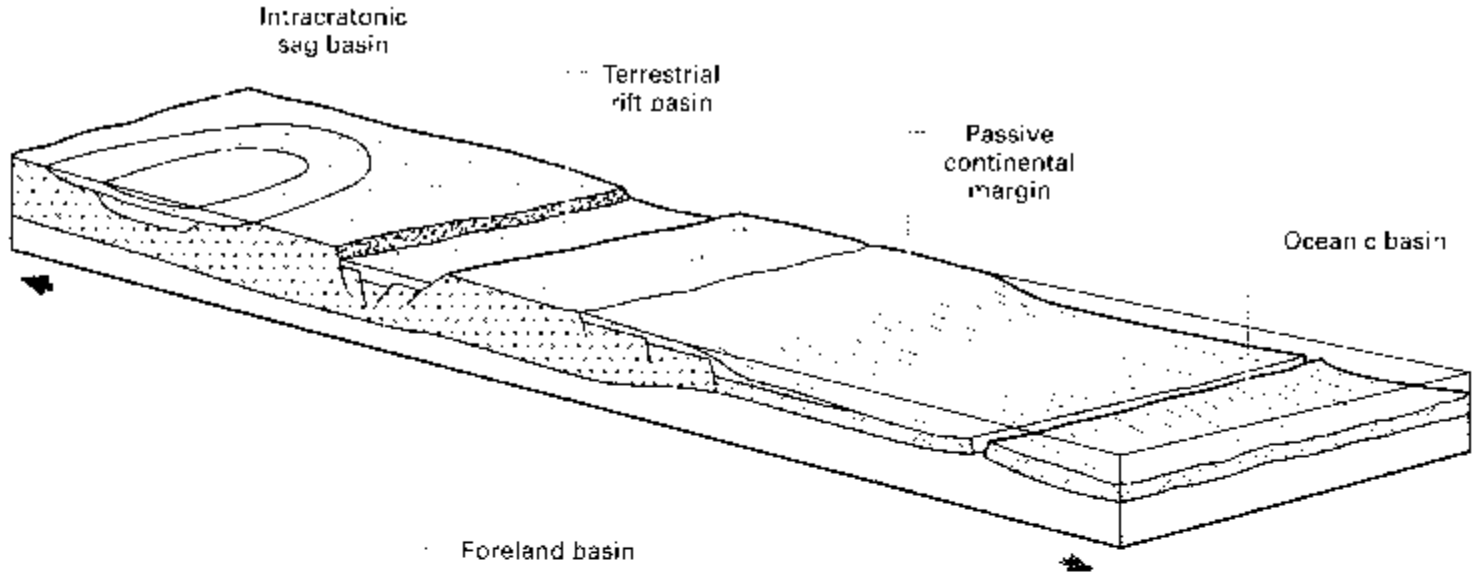


GEOL-201

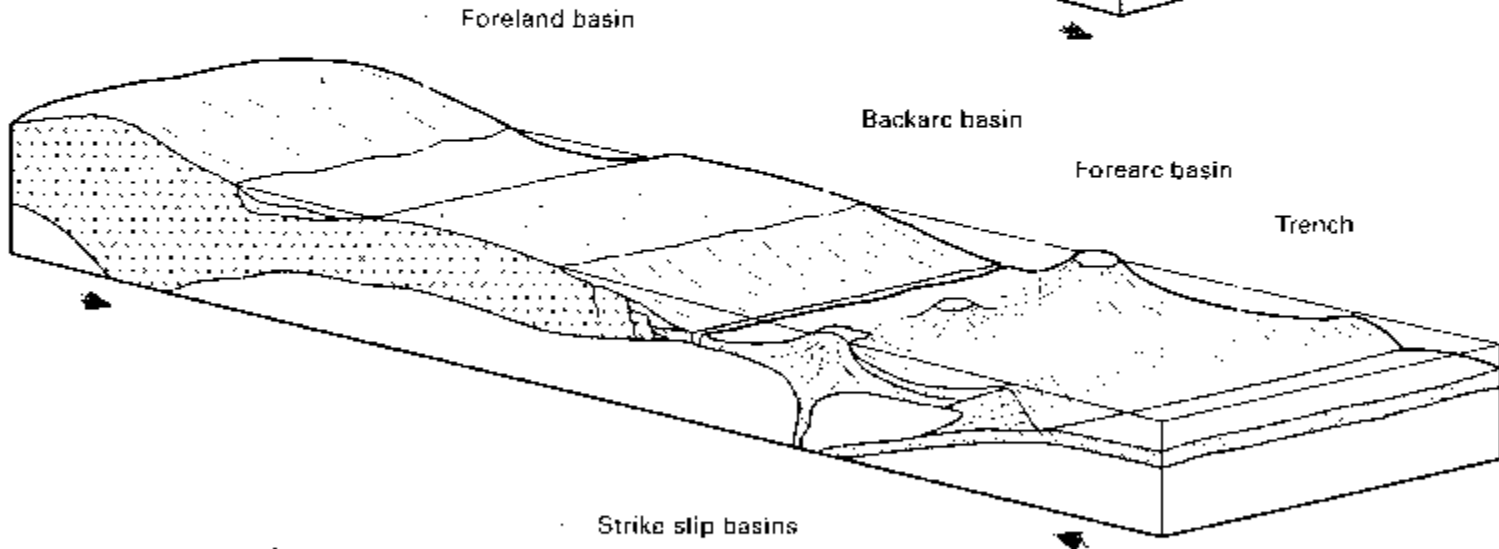
Sedimentary Basins



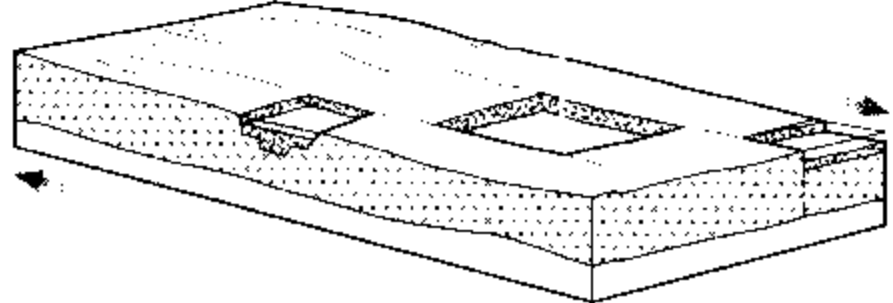
Cratonic basins

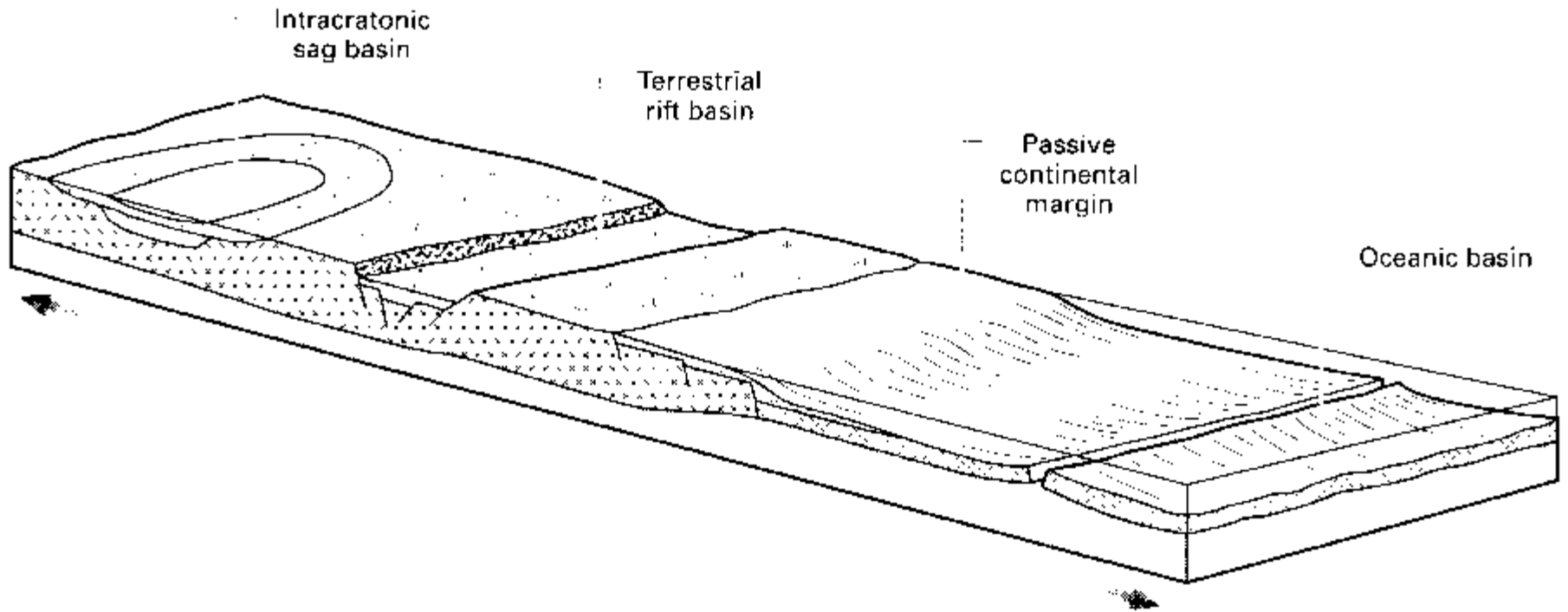


Convergence related basins



Strike-slip basins



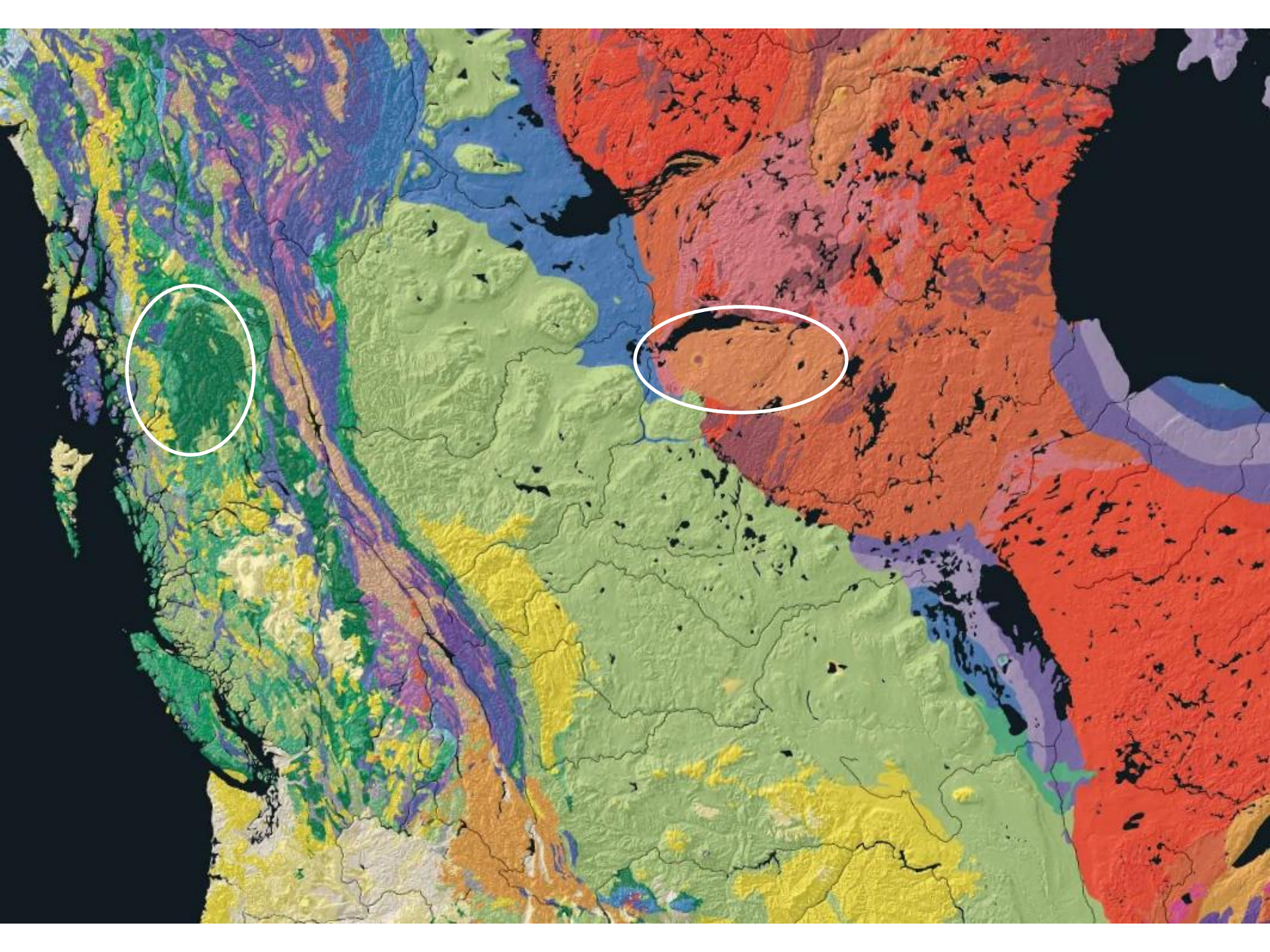


Intracratonic basins form within stable continental interiors. A good example is the Western Canada Sedimentary Basin extending from the eastern side of the Rocky Mountains to central Manitoba.

Examples of a smaller intracratonic basins are the Proterozoic Athabasca Basin in Saskatchewan and the Mesozoic Bowser Basin in northwestern BC.

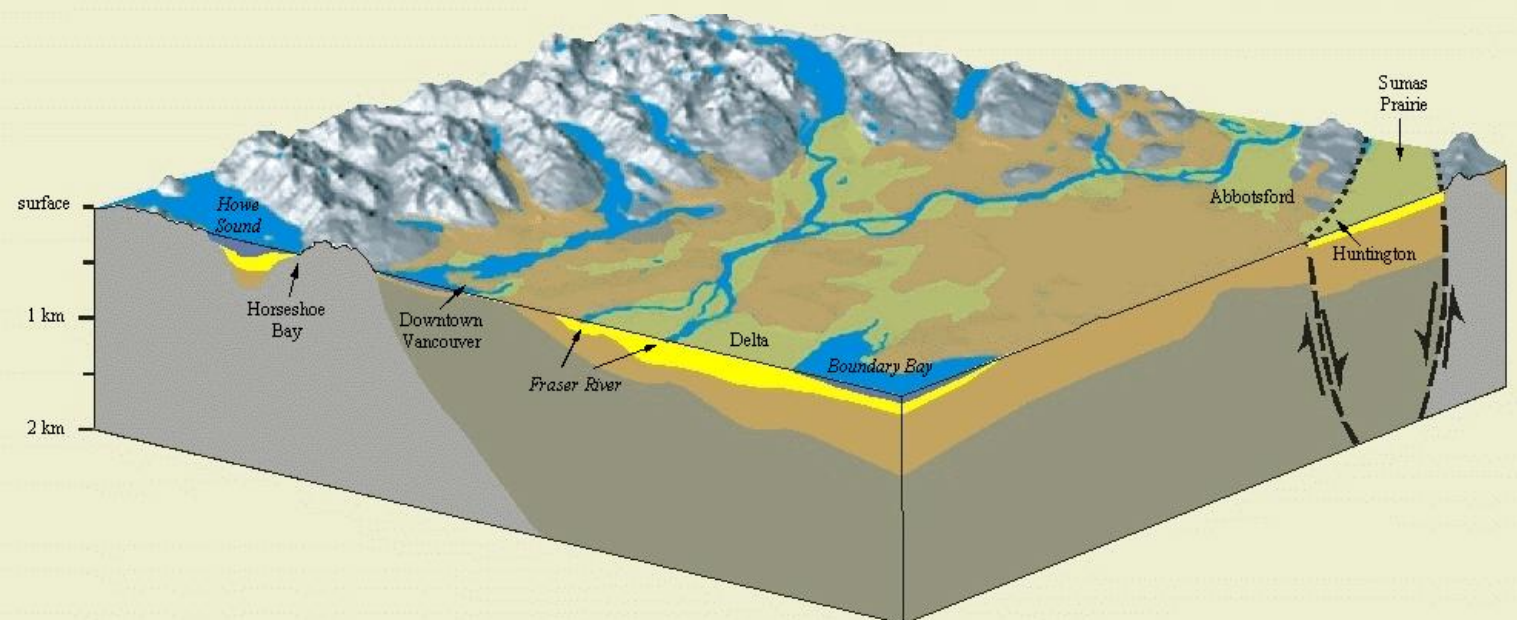
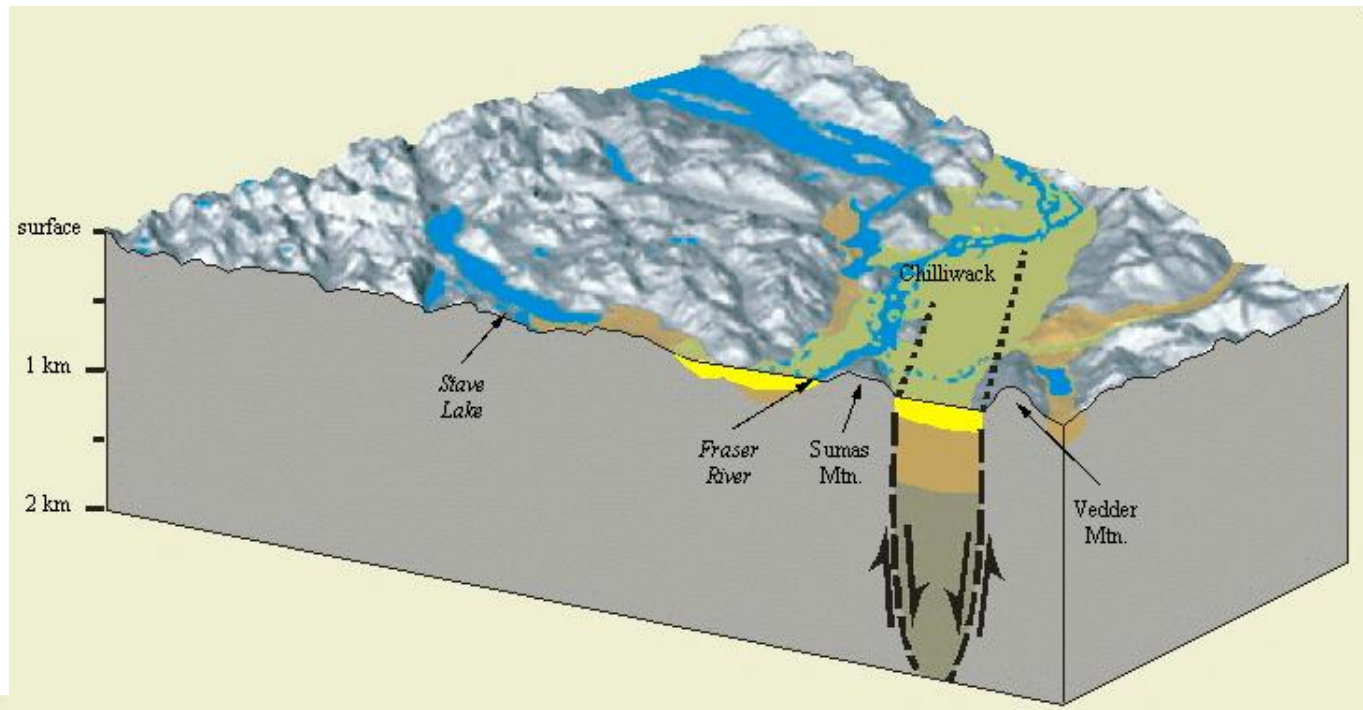
A modern example is Hudson Bay.

In all cases the underlying crustal rock is continental not oceanic.



Terrestrial rift basins form during rifting of a continent. The most obvious modern example is the East African rift. Rift basins don't necessarily result in continental rifting.

Other modern examples are Lake Baikal in Russia and the valley of the Rhine River in Europe (Rhine Graben) and the eastern part of the Fraser Valley

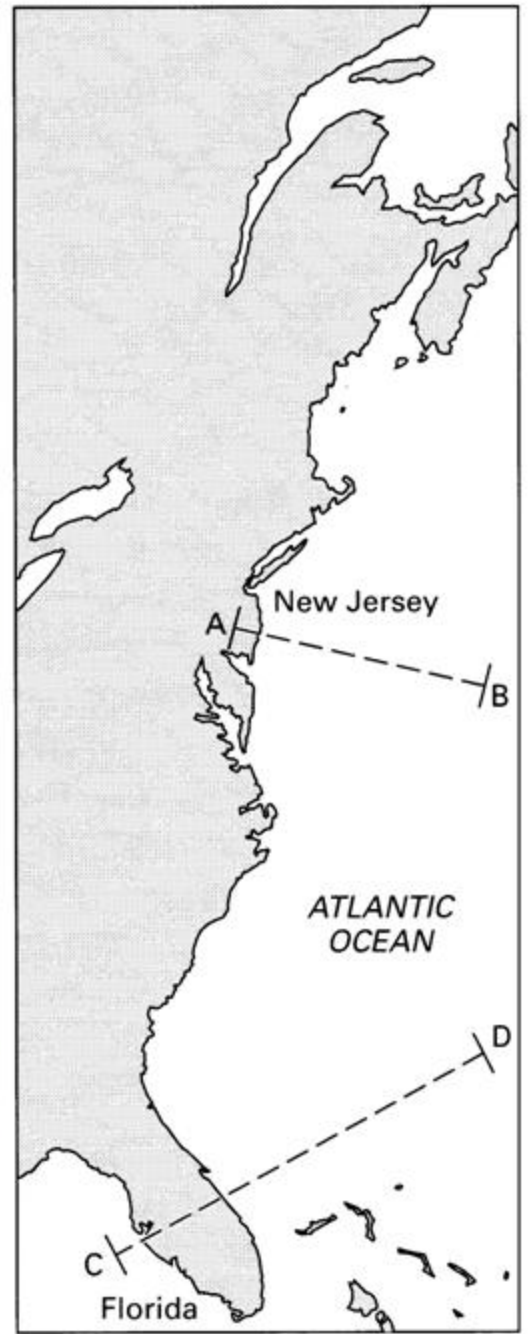
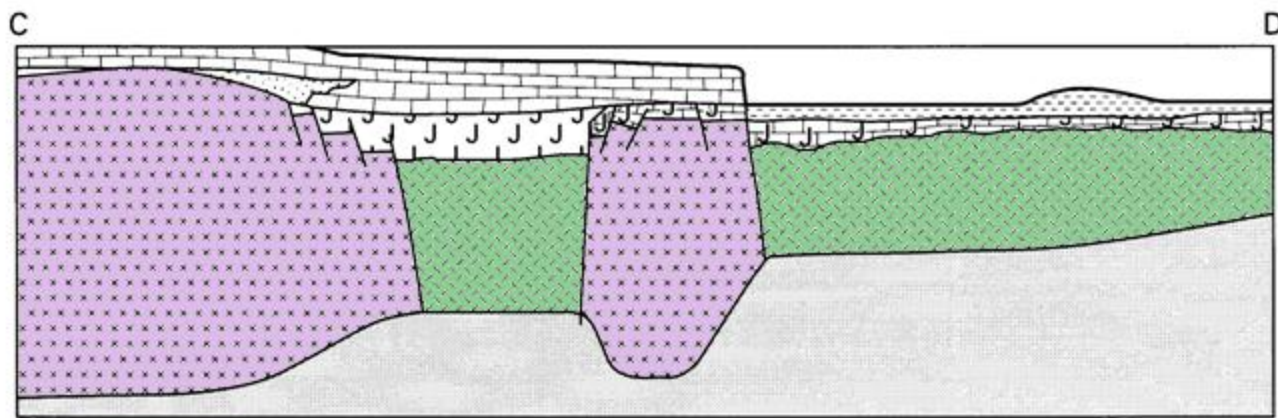
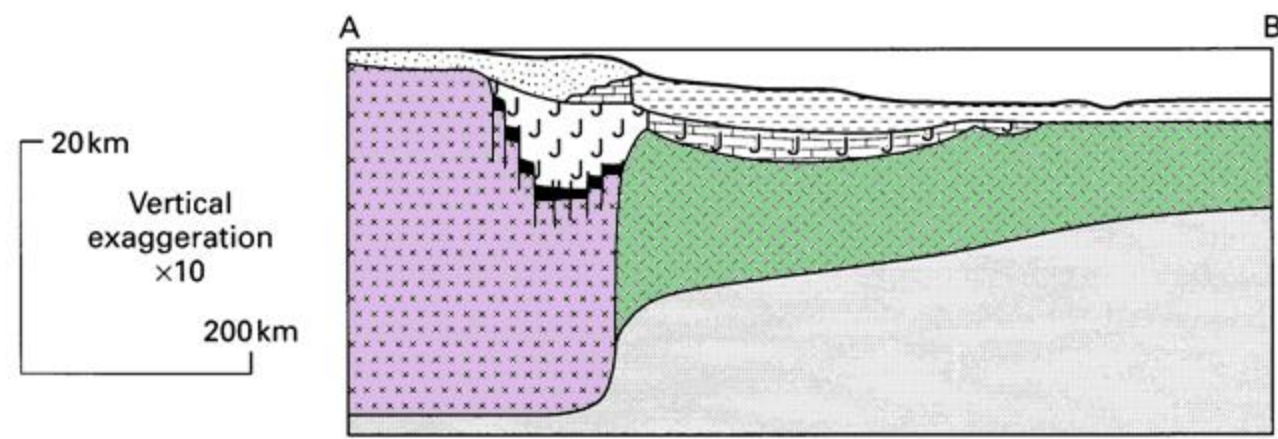
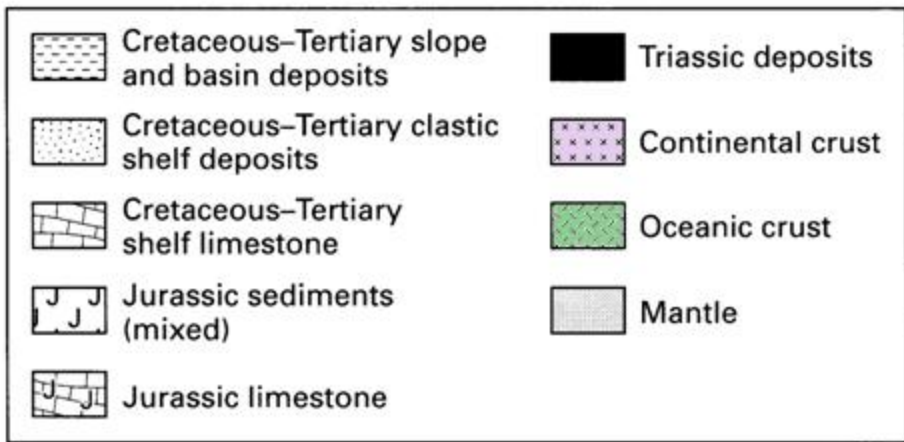


From GeoMap Vancouver

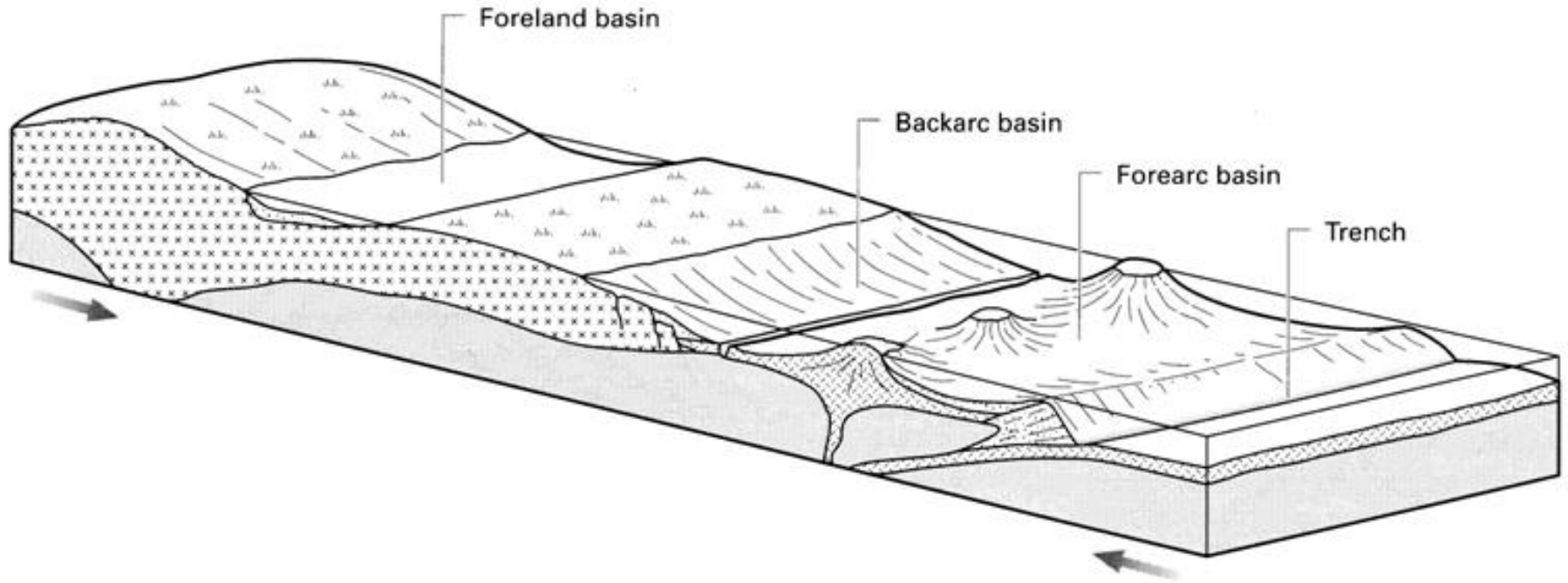
Passive margin basins form along continental-ocean margins where no subduction is taking place. Examples include the eastern coasts of North and South America and the US Gulf Coast.

The west coast of North America was a passive margin basin from around 700 to 200 Ma and the sediments are now seen in the Rocky Mts.

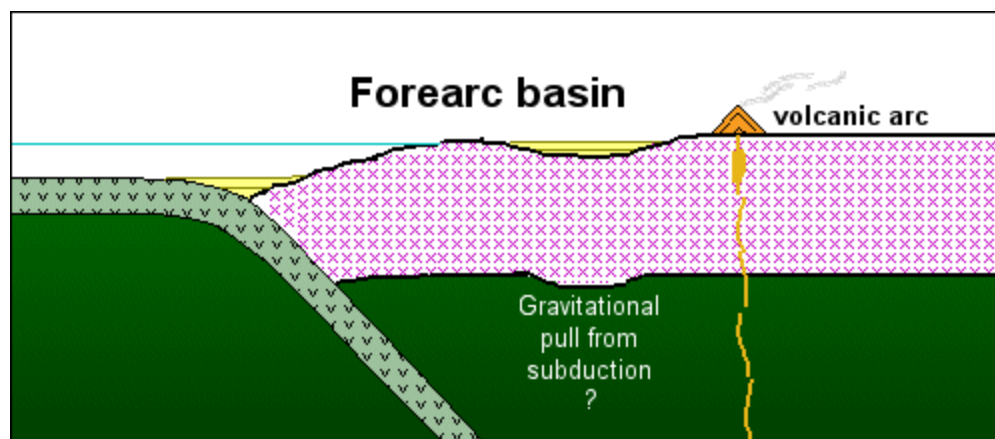
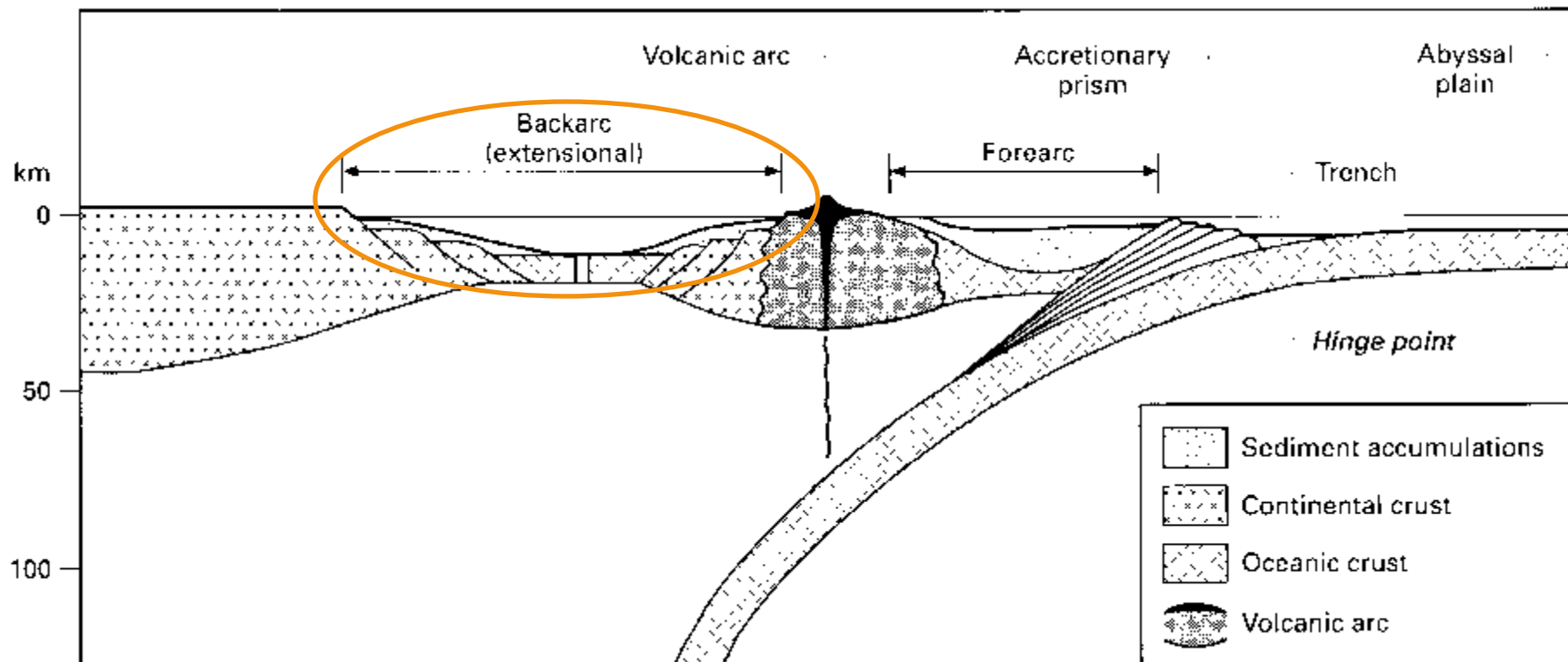
PMBs can accommodate sediment accumulations as much as 20 km thick.



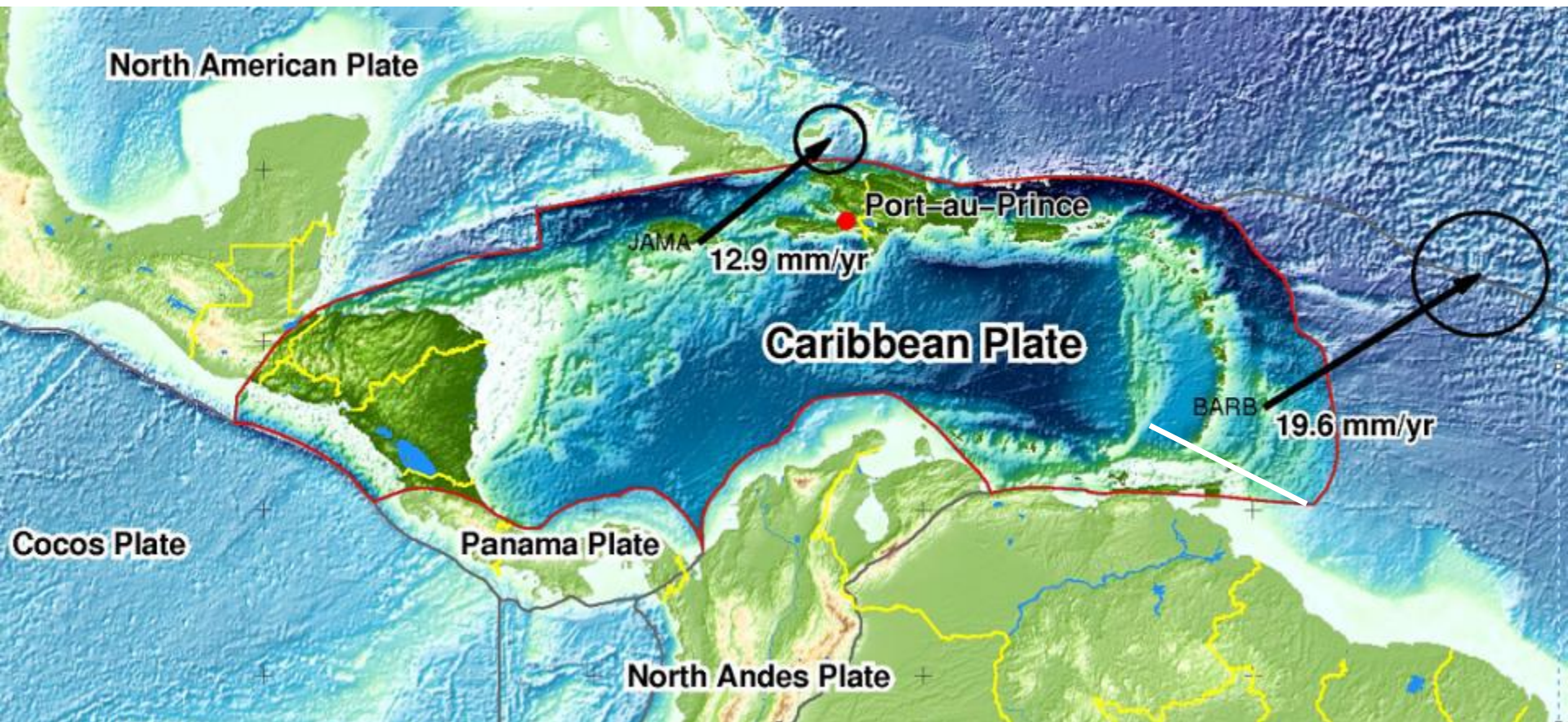
Convergence basins

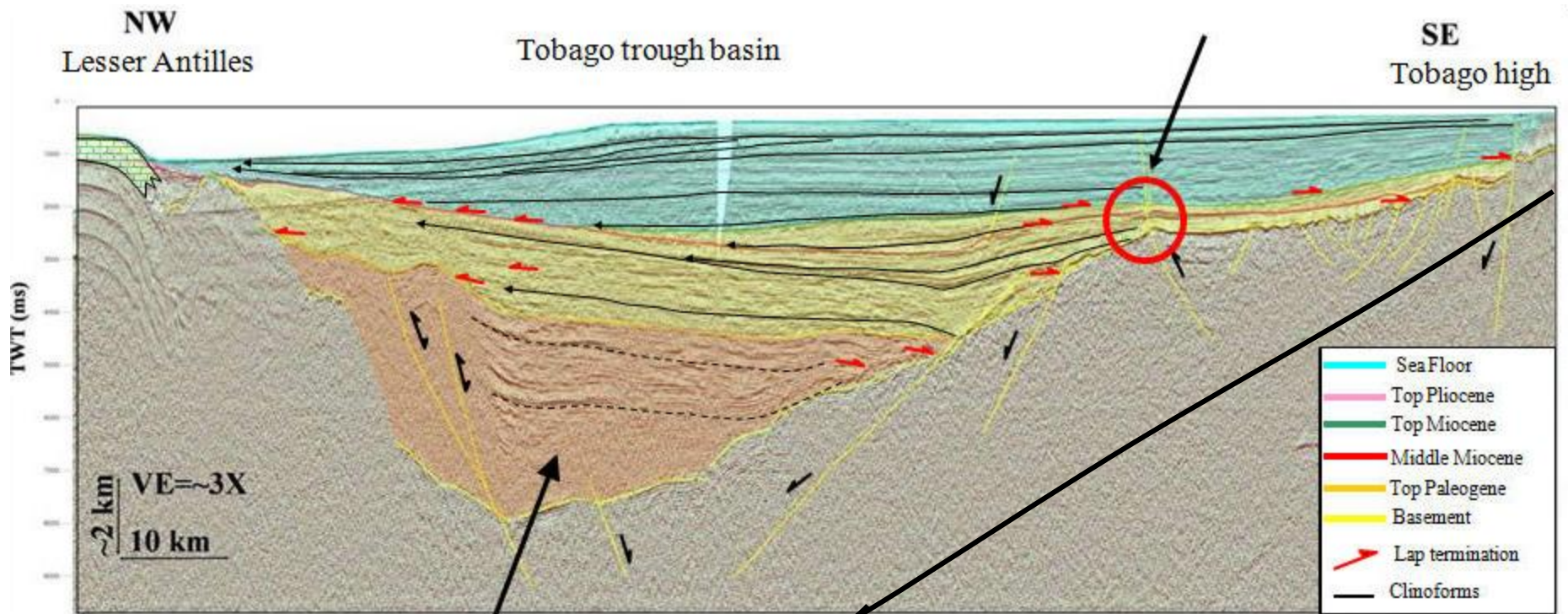


Trench, forearc, foreland and backarc basins form at active subduction boundaries.



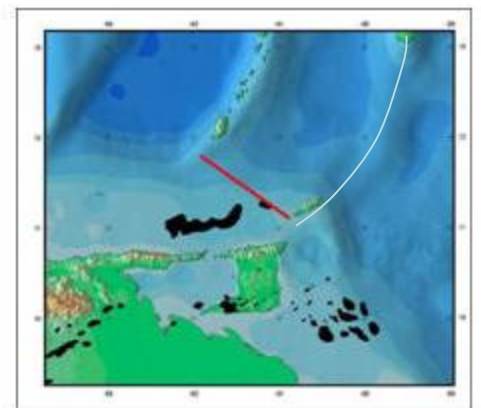






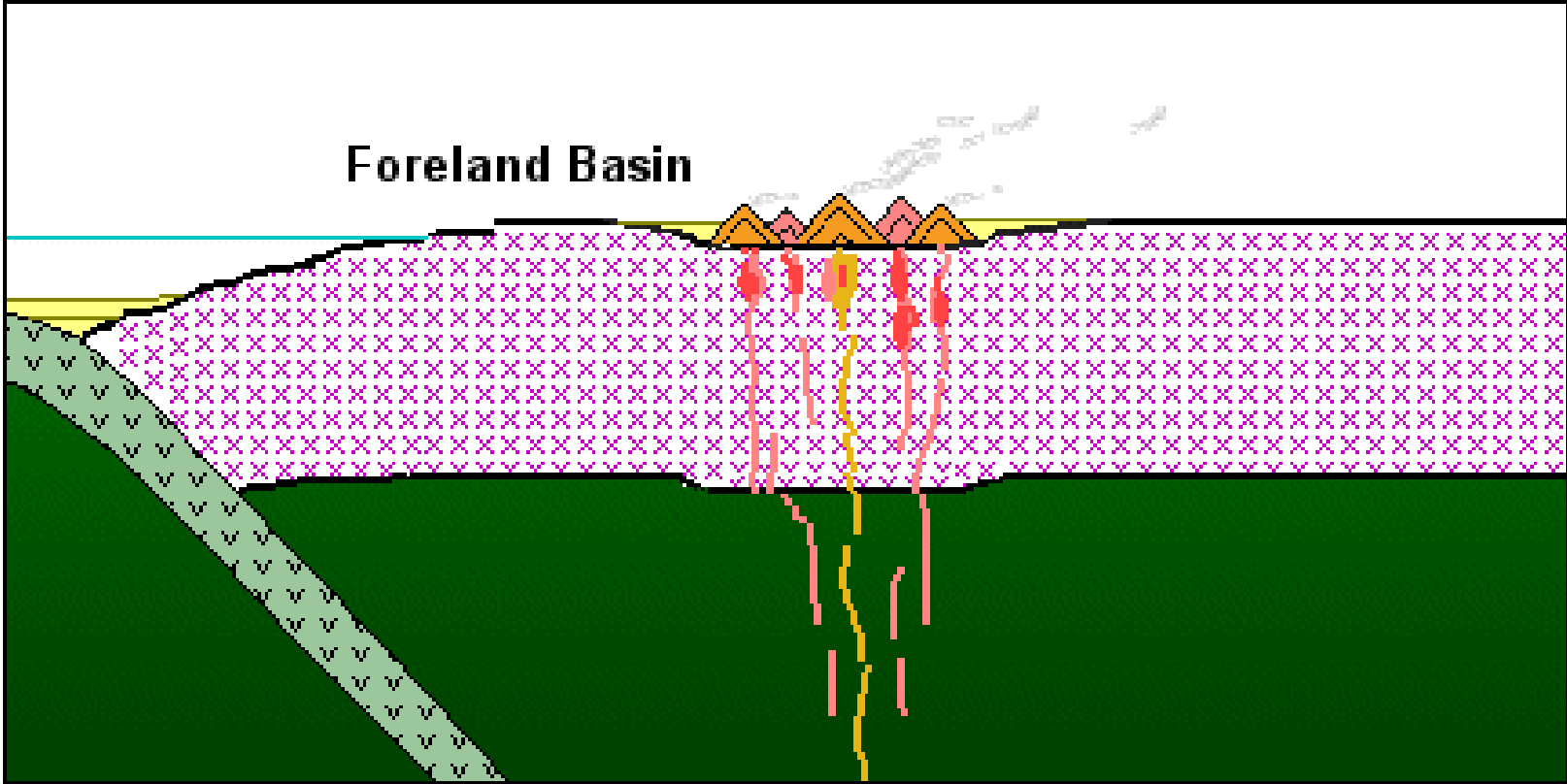
Paleogene: Potential source and reservoir rocks

None or little indication of Cretaceous source rocks in this part of the Tobago basin

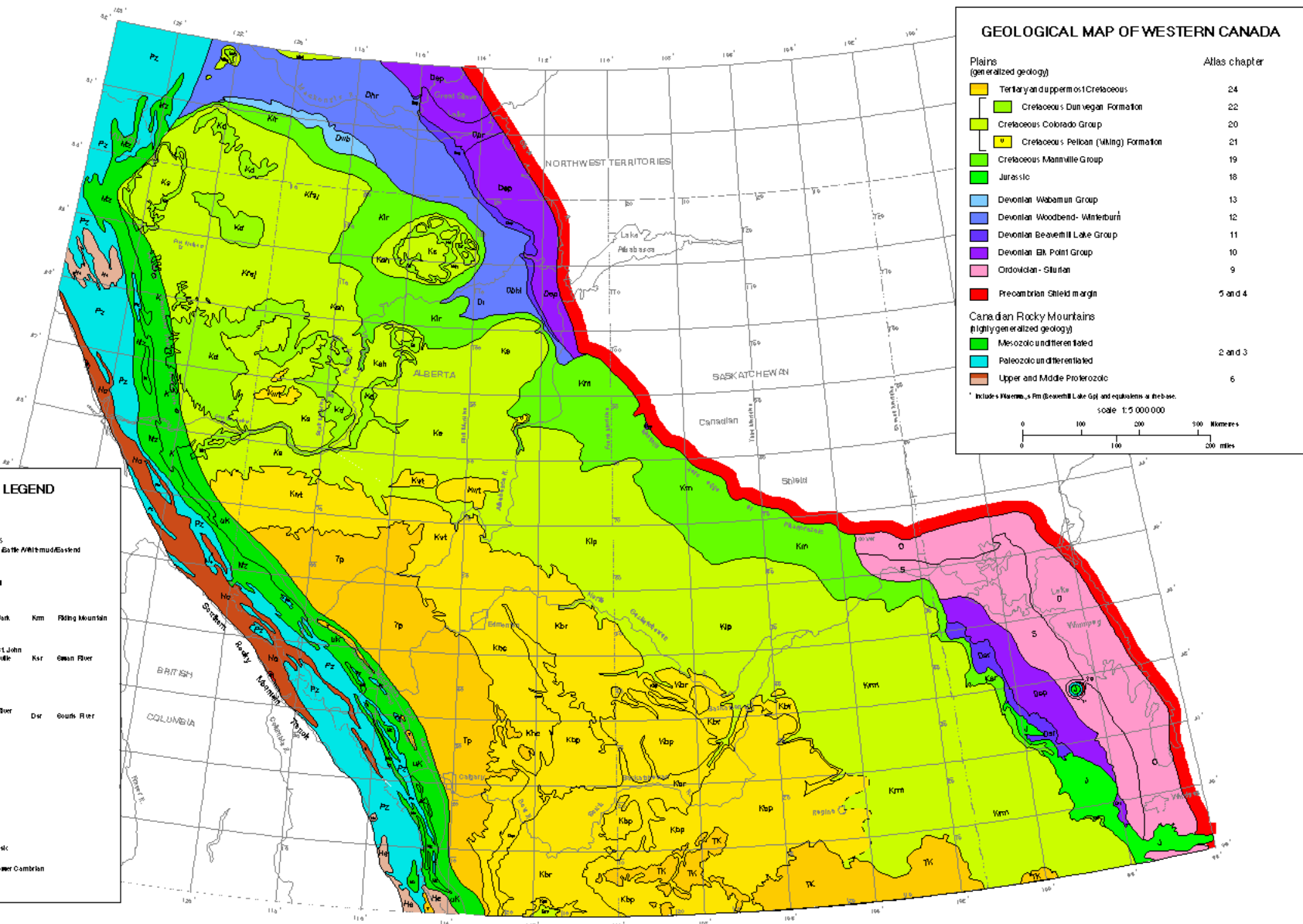


Foreland basins exist at convergent boundaries where large mountain belts have formed, such as the Coast Range, Andes, Rockies, Alps and Himalayas.

Foreland Basin



The WCSB is an intracratonic foreland basin

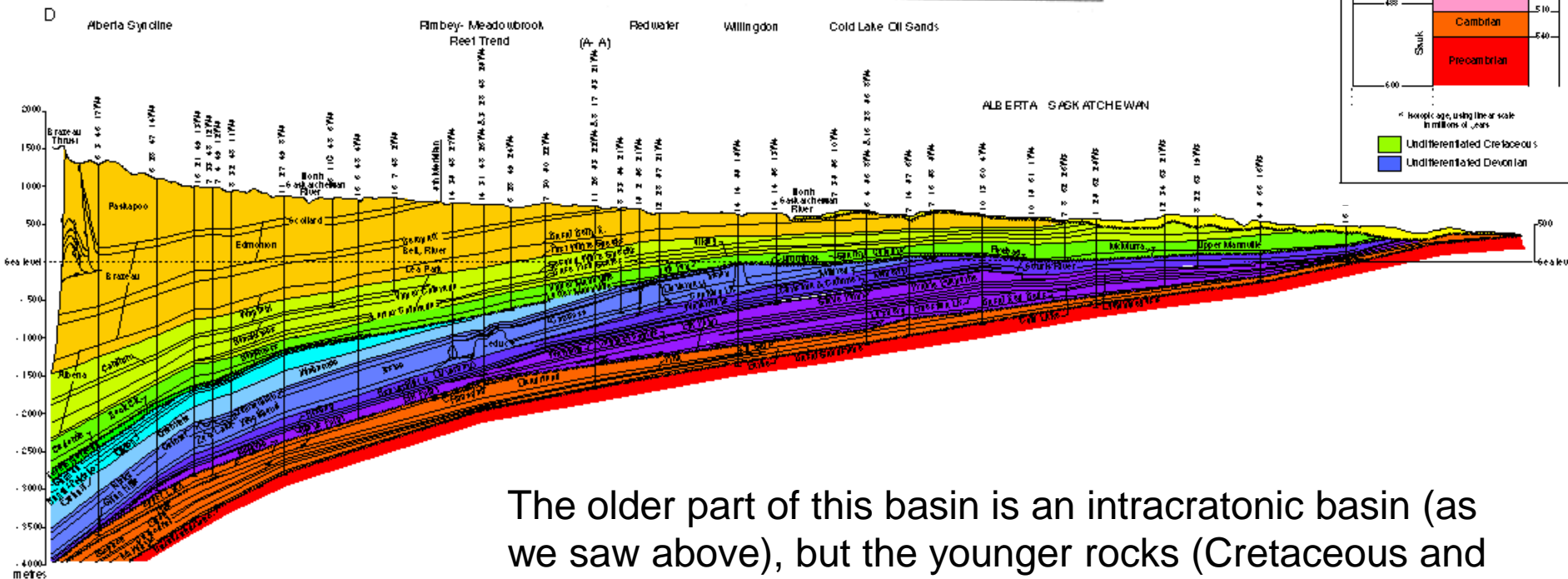


WCSB (Foreland Basin)



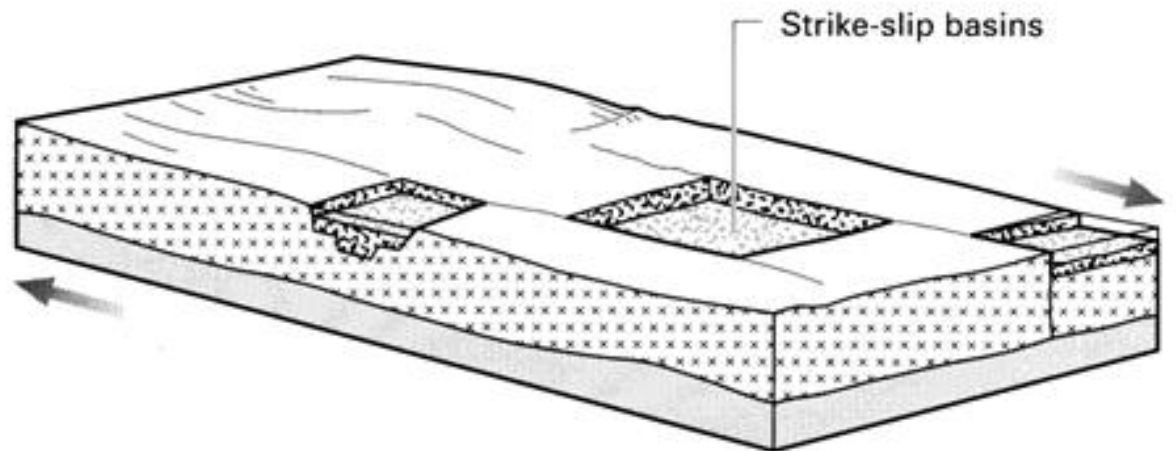
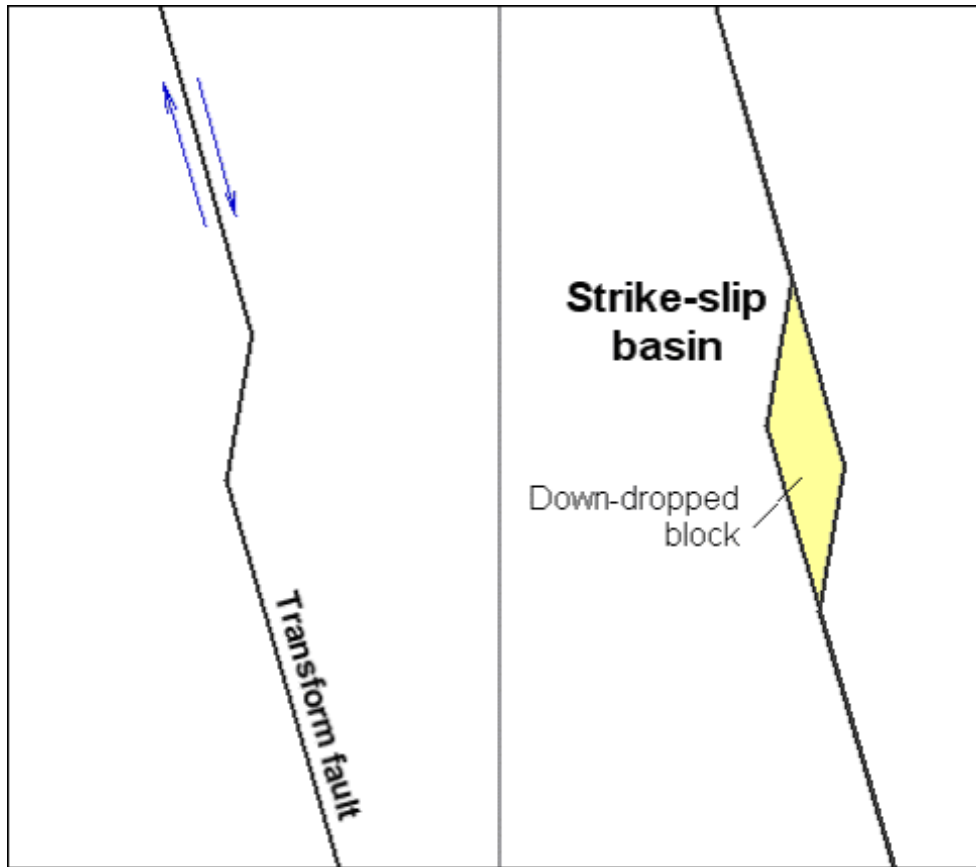
Geological Time Chart		
Sequences	Systems	
Tertiary	Quaternary	0 - c
	Tertiary and uppermost Cretaceous	0 - 65
Zuni	Cretaceous	65 - 146
	Jurassic	146 - 186
Algonkian	Triassic	186 - 208
	Permian	208 - 252
	Pennsylvanian	252 - 290
	Mississippian	290 - 330
Kaskaskian	Devonian	330 - 360
	Silurian	360 - 401
Tippesewan	Ordovician	401 - 443
	Cambrian	443 - 510
Shukon	Cambrian	510 - 540
	Precambrian	540 - 600

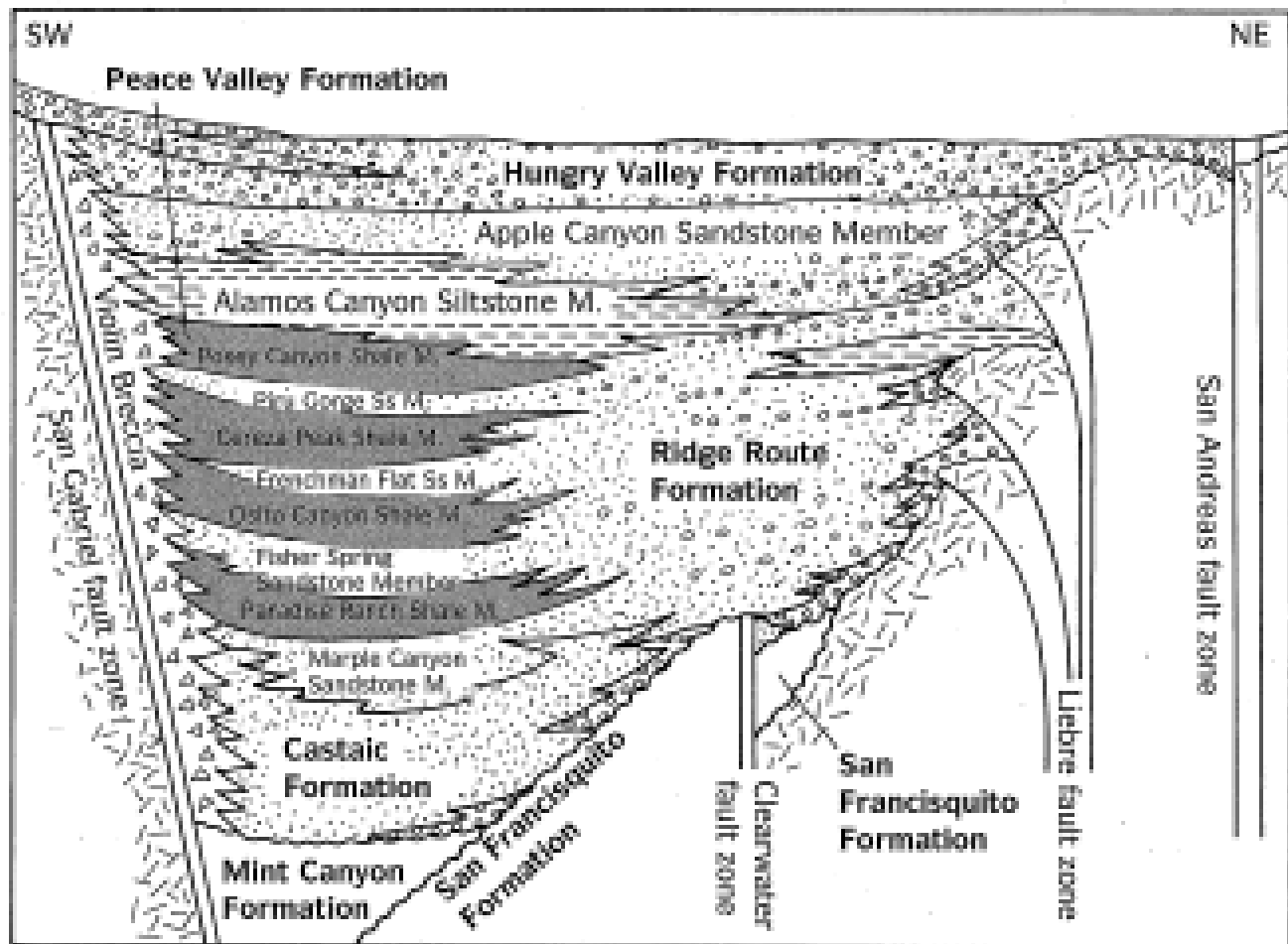
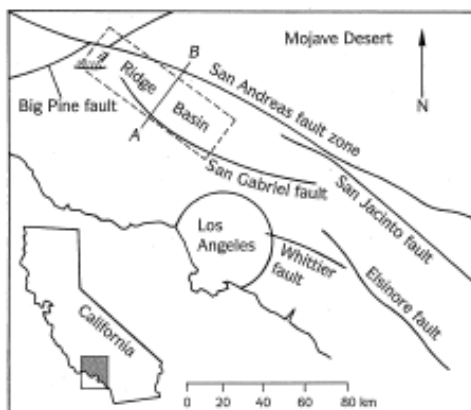
* Isotope age, using the scale in millions of years
 Legend:
 Undifferentiated Cretaceous (Yellow)
 Undifferentiated Devonian (Blue)



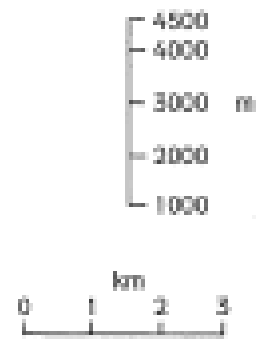
The older part of this basin is an intracratonic basin (as we saw above), but the younger rocks (Cretaceous and Devonian) accumulated in a foreland basin.

Pull-apart or strike-slip basins form along major transform faults where they bend, or between fault splays. They tend to be deep and narrow, and are characterized by very fast subsidence. There are some good examples along the San Andreas fault-zone.





-  Breccia
-  Conglomerate
-  Sandstone
-  Siltstone
-  Shale (mudstone)
-  Basement rocks





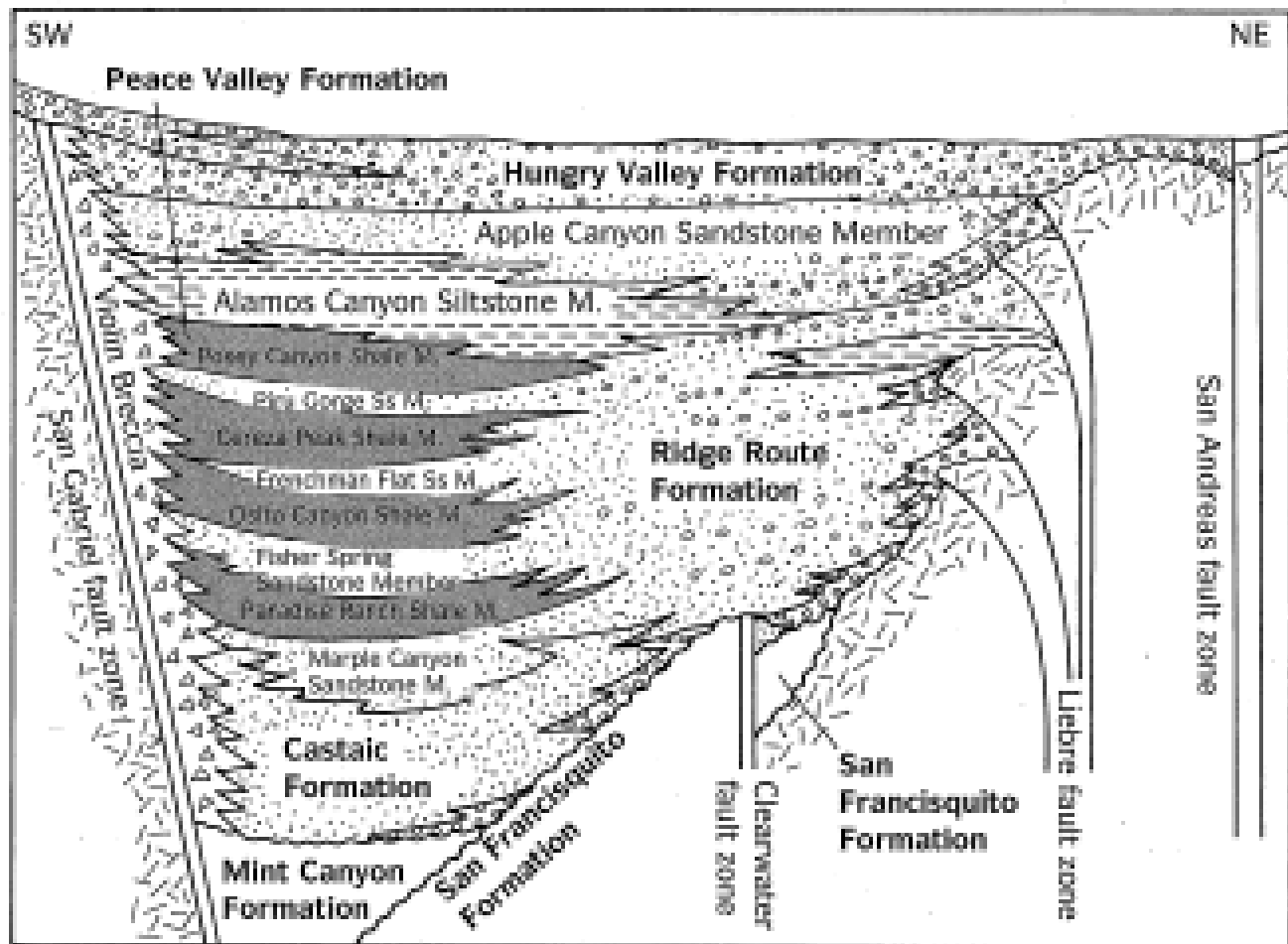
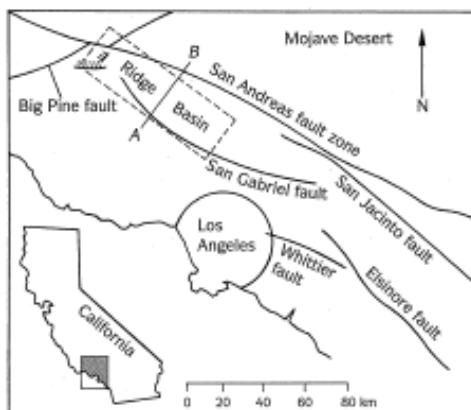
Accommodation

In the context of sedimentary basins “accommodation” refers to the amount of vertical space available for sediments to accumulate.

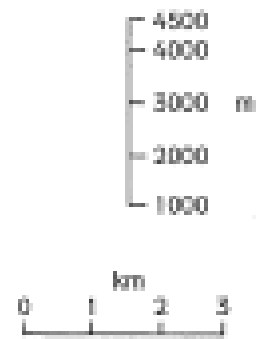
Accommodation

When sediments accumulate, the basin in which they are being deposited will tend to subside because of isostasy.

If a basin is filling up faster than the rate of subsidence then the deposition will move elsewhere – either to a new basin – or farther out into the ocean.



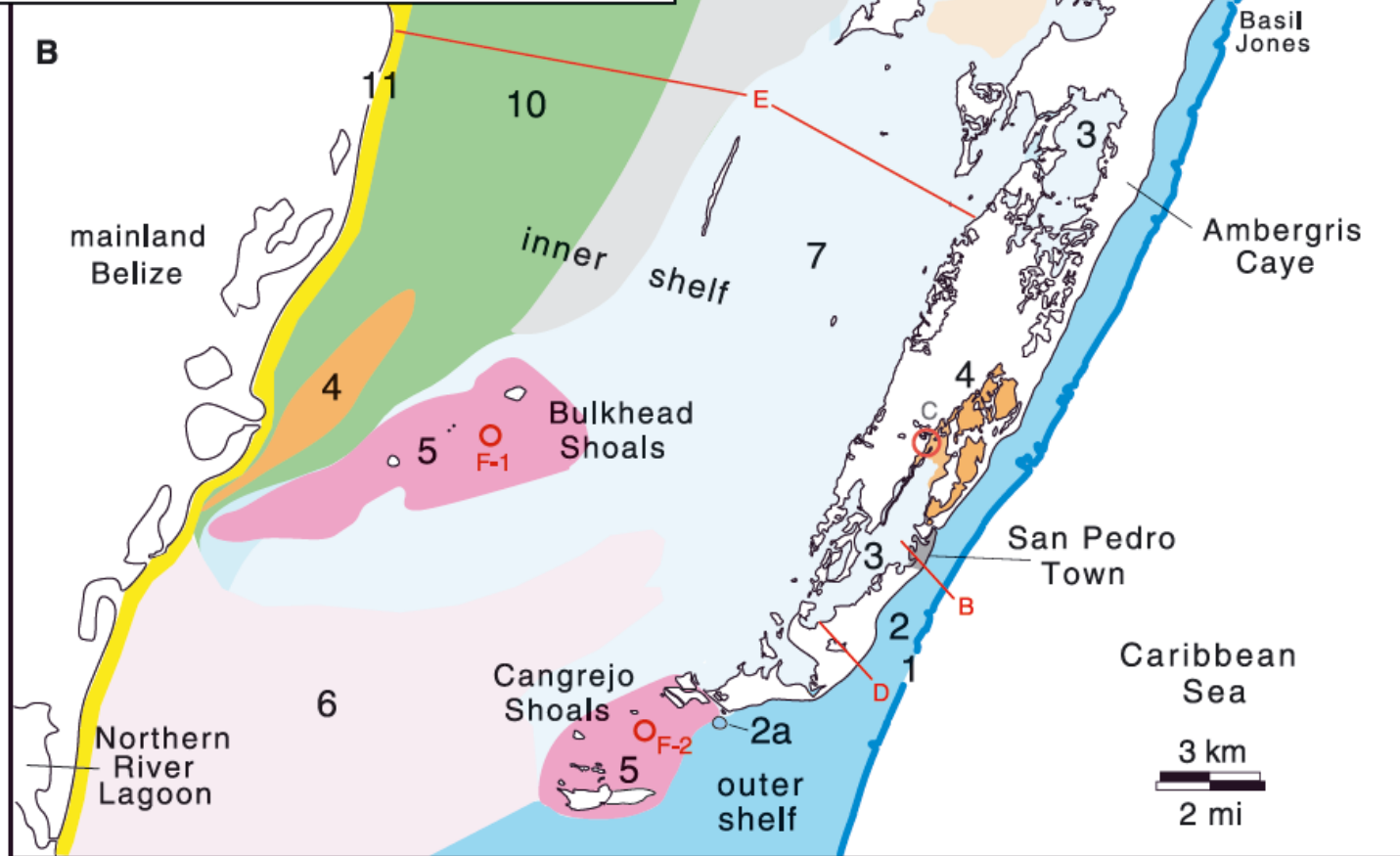
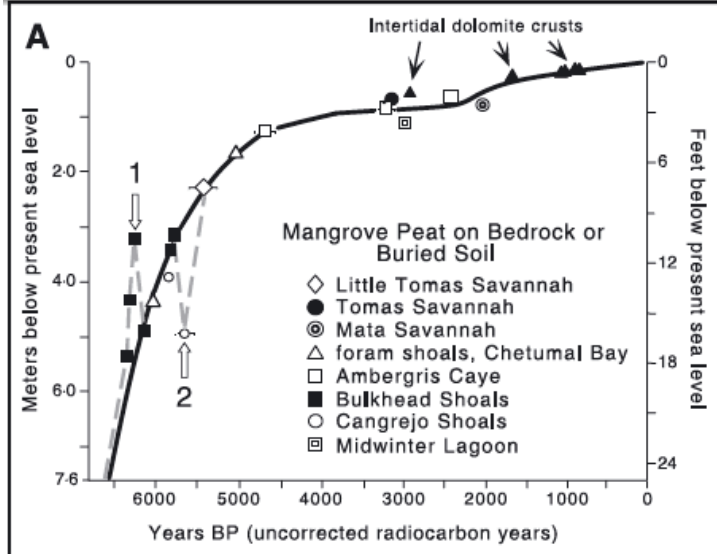
-  Breccia
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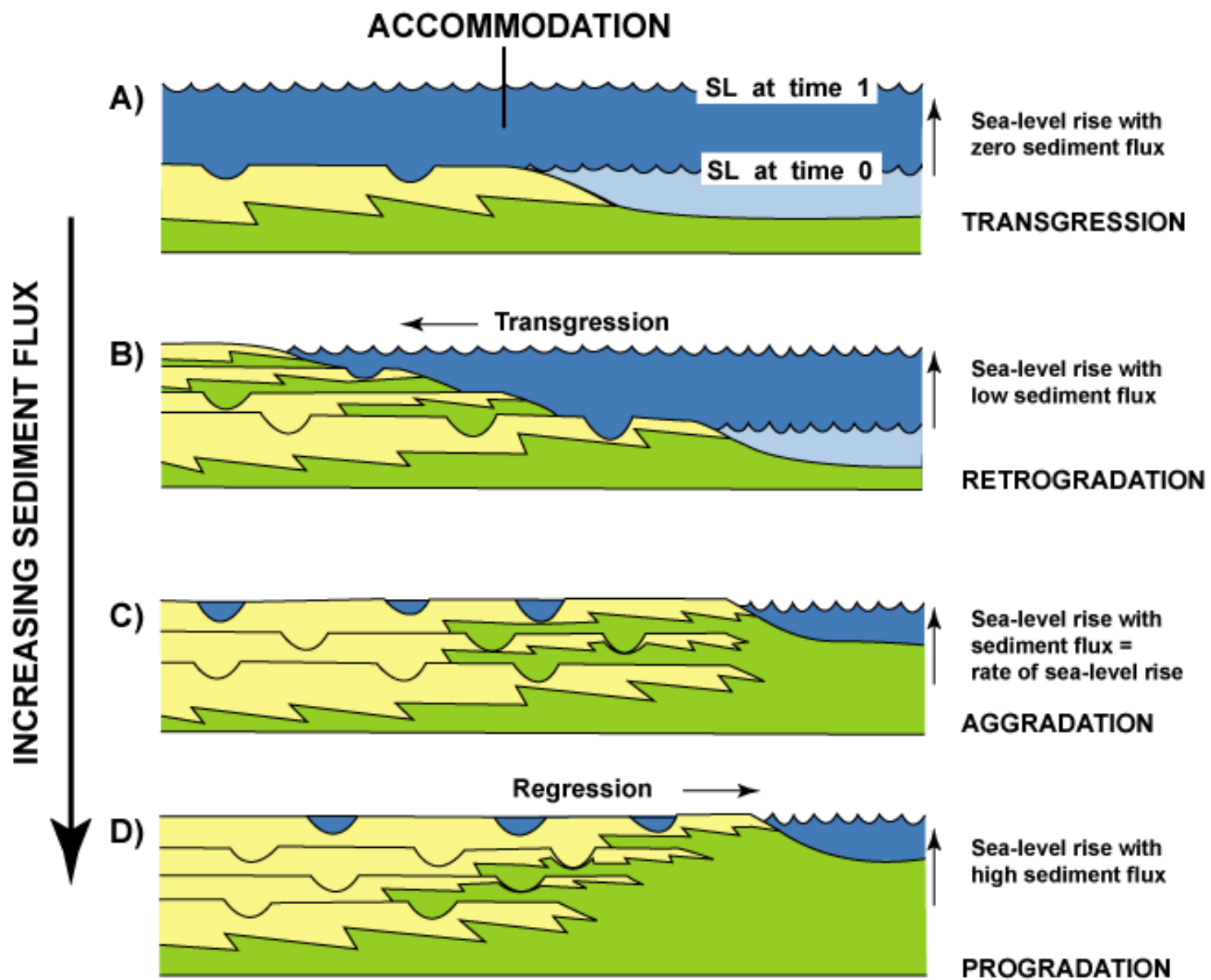


Accommodation

Accommodation is affected by other tectonic processes and by sea-level rise.

Sea level rose dramatically after the end of the last glaciation and the impact is observed on shallow marine shelves like the one in Belize.





After Posamentier & Allen, 1999